

Mathematisch-Naturwissenschaftliche Fakultät

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Suggested citation referring to the original publication: Biogeosciences 14 (2017), pp. 1617–1630 DOI https://doi.org/10.5194/bg-14-1617-2017 ISSN (print) 1726-4170 ISSN (online) 1726-4189

Postprint archived at the Institutional Repository of the Potsdam University in: Postprints der Universität Potsdam Mathematisch-Naturwissenschaftliche Reihe ; 672 ISSN 1866-8372 https://nbn-resolving.org/urn:nbn:de:kobv:517-opus4-417104 DOI https://doi.org/10.25932/publishup-41710

Biogeosciences, 14, 1617–1630, 2017 www.biogeosciences.net/14/1617/2017/ doi:10.5194/bg-14-1617-2017 © Author(s) 2017. CC Attribution 3.0 License.



C / N ratio, stable isotope (δ^{13} C, δ^{15} N), and *n*-alkane patterns of brown mosses along hydrological gradients of low-centred polygons of the Siberian Arctic

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Received: 13 April 2016 – Discussion started: 29 July 2016 Revised: 7 February 2017 – Accepted: 16 February 2017 – Published: 28 March 2017

Abstract. Mosses are a major component of the arctic vegetation, particularly in wetlands. We present C / N atomic ratio, δ^{13} C and δ^{15} N data of 400 brown-moss samples belonging to 10 species that were collected along hydrological gradients within polygonal mires located on the southern Taymyr Peninsula and the Lena River delta in northern Siberia. Additionally, n-alkane patterns of six of these species (16 samples) were investigated. The aim of the study is to see whether the inter- and intraspecific differences in C/N, isotopic compositions and *n*-alkanes are indicative of habitat, particularly with respect to water level. Overall, we find high variability in all investigated parameters for two different moisture-related groups of moss species. The C/N ratios range between 11 and 53 (median: 32) and show large variations at the intraspecific level. However, species preferring a dry habitat (xero-mesophilic mosses) show higher C / N ratios than those preferring a wet habitat (meso-hygrophilic mosses). The δ^{13} C values range between -37.0 and -22.5% (median = -27.8%). The δ^{15} N values range between -6.6 and +1.7% (median = -2.2%). We find differences in δ^{13} C and δ^{15} N compositions between both habitat types. For some species of the meso-hygrophilic group, we suggest that a relationship between the individual habitat water level and isotopic composition can be inferred as a function of microbial symbiosis. The *n*-alkane distribution also shows differences primarily between xeromesophilic and meso-hygrophilic mosses, i.e. having a dominance of *n*-alkanes with long $(n-C_{29}, n-C_{31})$ and intermediate $(n-C_{25})$ chain lengths, respectively. Overall, our results reveal that C/N ratios, isotopic signals and *n*-alkanes of studied brown-moss taxa from polygonal wetlands are characteristic of their habitat.

1 Introduction

Specific physiological and morphological traits enable mosses to attain extensive surface coverage in lowtemperature ecosystems such as arctic tundra (Turetsky et al., 2012; Wasley et al., 2006). Mosses are a major component of the biomass in arctic wetlands and contribute strongly to the biodiversity. In particular, the vegetation of the widely distributed polygonal tundra is rich in moss taxa, which partly originates from the strong compositional turnover along a small-scale hydrologic gradient (Zibulski et al., 2016). As an intermediate layer between air and the permafrost soil, mosses control the water cycle, greenhouse gas and energy exchange (Blok et al., 2011; McFadden et al., 2003), as well as the structure of the habitats of vascular plant communities in arctic landscapes (Beringer et al., 2001; Gornall et al., 2011; Malmer et al., 1994). Because of their low decomposition rates (Aerts et al., 1999; Turetsky, 2003), they contribute strongly to the peat and permafrost carbon pool (Tarnocai et al., 2009).

Despite the significance of mosses in high-latitude biodiversity and matter cycles, little is known about their carbonto-nitrogen ratio (C/N), stable isotope and *n*-alkane characteristics in comparison to vascular plants. Such information is not only necessary to improve our understanding of the physiological adaptation or plasticity of mosses to certain environmental characteristics, but can also be used when similar measurements of fossil plant material are interpreted as proxies of former ecological or environmental conditions (e.g. Birks, 1982).

C / N ratios, stable carbon and nitrogen isotope values $(\delta^{13}C, \delta^{15}N)$, and the *n*-alkane fingerprints of bulk organic matter are among the most commonly measured parameters of plant matter (Chambers and Charman, 2004). The C/N ratios of mosses are generally in the range of those of higher terrestrial plants (Prahl et al., 1980) but are known to depend on the available nitrogen, which originates in pristine regions from decomposition of organic matter, microbial activity or atmosphere deposits (e.g. Chanway et al., 2014; Lee et al., 2009). We provide C / N ratios by weight of arctic brown mosses, in anticipation that they will be useful for comparative palaeo-environmental reconstructions (Andersson et al., 2011) and in the evaluation of organic matter sources in Russian permafrost soils with regard to species and habitatspecific patterns. Furthermore, C / N ratios may be related to growth form, i.e. higher ratios can be expected for mosses which compete with vascular plants for light and thus need to invest in a high stem stability (Sveinbjörnsson and Oechel, 1992).

As well as the C / N ratio, the δ^{13} C composition is barely altered by decomposition processes in the soil and is also suitable for palaeo-environmental reconstructions. The δ^{13} C ratio of an individual plant are a mixed signal of the physiological traits of the species and the individual's direct environment. Mosses use the C₃ pathway for carbon assimilation (O'Leary, 1988; Farquhar et al., 1989) and despite a lack of stomata in the photosynthetically active parts, they have a similar range in their δ^{13} C values of between -24and -32% (Ménot and Burns, 2001; Rundel et al., 1979; Smith and Epstein, 1971) as vascular C3 plants. Differences in δ^{13} C values among several species can be explained by individual plant physiology and biochemistry (Galimov, 2000). Differences within a single species have been linked with environmental conditions such as temperature (Skrzypek et al., 2007; Waite and Sack, 2011), growing depth below water level (Raghoebarsing et al., 2005), position within a cushion or hummock (Price et al., 1997), lipid content (Rundel et al., 1979), or the influence of microbial symbioses (Liebner et al., 2011; Vile et al., 2014). Furthermore, even differences among branches and stems of single individuals have been reported (Loader et al., 2007). However, almost all of these studies were made on Sphagnaceae, which are not representative of all mosses because of their specific morphology (i.e. the occurrence of a photosynthetically active cell type and a dead cell type, which is responsible for water storage and often an additional coating of the photosynthetic active cell) and their specific habitat preferences (i.e. preferring acidic wetlands). Hence, this information cannot simply be transferred to brown mosses – which form the major component in northern Siberian lowlands – due to their different morphology.

Most studies on δ^{15} N values of moss material have focused on the anthropogenic impact on the nitrogen cycle (Harmens et al., 2011; Liu et al., 2008; Poikolainen et al., 2009), whereas reports on the δ^{15} N compositions of mosses from relatively pristine ecosystems such as the Arctic are rare or have been investigated in relation to the study of bird colonies (e.g. Lee et al., 2009). Potentially, such information can indicate pathways and sources of the nutrient supply in these N-limited ecosystems (Kielland, 1997; Michelsen et al., 1996, 1998).

Compared with vascular plants and Sphagnum species, relatively few investigations of *n*-alkane patterns of brown mosses are available. Palaeo-environmental reconstructions use the potential of *n*-alkanes to distinguish between different plant groups (Ficken et al., 1998), between moisture conditions (Pancost et al., 2000; Nichols et al., 2006; Zhou et al., 2010), whether organic material is to decomposed, and between changes in temperature (Feakins et al., 2016; Sachse et al., 2006) along distinct gradients. Sphagnum species, for example, show a dominance of $n-C_{23}$ and $n-C_{25}$ homologues (Baas et al., 2000) that are comparable to the pattern of vascular submerged plants (Ficken et al., 2000), and ratios are used as a proxy for wet moisture conditions. Other studies discuss the suitability of *n*-alkane patterns in moss species for chemotaxonomical studies on recent and fossil material (Bush and McInerney, 2013; Nott et al., 2000; Schellekens and Buurman, 2011). A greater protection potential of waxes with a higher content of long-chain n-alkanes against solar irradiation or, alternatively, an enhanced loss of shortchain *n*-alkanes by evaporation have been suggested as possible mechanisms to cause different *n*-alkane patterns in leaf waxes of trees (Sachse et al., 2006). However, with respect to mosses the pattern and mechanisms are even less understood.

This study presents C / N, isotopic (δ^{13} C, δ^{15} N) and *n*alkane characteristics of mosses from low-centred polygons in northern Siberia. Low-centred polygons are geomorphological forms in arctic landscapes originating from frostheave processes in the soil. They are characterized by elevated dry rims and a water-saturated or water-filled centre. This centimetre-scale hydrological gradient is well reflected by a strong turnover in the vascular plant and moss composition (Zibulski et al., 2016). We investigate the relationship between the C / N, isotopic and *n*-alkane compositions and the hydrological conditions within low-centred polygons. In particular, we aim to reveal whether fingerprints are habitatspecific or rather species-specific.

2 Material and methods

2.1 Sites

The plant material was collected during the vegetation season (July-August) from eight low-centred polygons located along a zonal vegetation gradient ranging from open forest via the forest-tundra intersection to subarctic tundra (Matveev, 1989) to obtain a representative sample set of northern Siberian lowlands (Fig. 1). Six of the polygons (06/P, 17/P, P3/I, P3/II, P3/III, 12/P), sampled in 2011, are located in the Khatanga River region (70–72° N, 97–102° E, southern Taymyr Peninsula) and a further two polygons (LP1 and LP2), sampled in 2012, are located on Samoylov Island in the Lena River Delta (72.4° N, 126.5° E). The climate is cold-humid (Khatanga climate station annual mean temperature and precipitation: -13.2 °C and 272 mm, Rivas-Martinez and Rivas-Saenz, 2009; Samoylov climate station annual mean temperature and precipitation: -12.5 °C and 232.7 mm, Boike et al., 2013).

2.2 Sampling and studied moss species

A representative continuous transect of adjacent 1 m^2 plots (from rim to rim) was selected for sampling in each polygon. The surface height in relation to water level (measured at the centre of each plot) and vegetation (abundance information) were recorded. Individual characteristics of each low-centred polygon are presented in Table S1 in the Supplement. Moss plants were hand-picked, dried in the field and transported to the Alfred Wegener Institute in Potsdam. Taxa identification is based on the relevant literature (Frahm and Frey, 2004; Lobin et al., 1995; Smith, 1978) and the Bryophyte Flora of North America internet source (http://www.efloras.org).

In total, eight species that were observed to have different water-regime preferences were included in the analyses. To approximate water-regime preferences of each species, we calculated the mean plant position in relation to water level (*h*) from the recorded plot surface height. *Warnstorfia exannulata* was observed to grow as part of a swinging mat at water level; it was assigned a preference for water level. Using this information the species were classified as either as of xero-mesophilic mosses (mean plant position between 13 and 16 cm in relation to water level) or to the group of meso-hygrophilic mosses (mean plant position between 3 and -30 cm in relation to water level) to ease the presentation of results (Table 1).



Figure 1. The Khatanga study sites are located in the south-east of the Taymyr Peninsula and both the sites on Samoylov Island are in the southern Lena River delta (Map by Th. Böhmer).

2.3 Measurements of C / N ratios, stable carbon and nitrogen isotope values and *n*-alkane distributions

Selected plant material (i.e. apical parts of a specimen) was rinsed with de-ionised water and mechanically cleaned from organic particulate material. The content of carbon and nitrogen and the ratio of stable isotopes were measured with a DELTAplusXL isotope ratio mass spectrometer (Thermo Fischer Scientific) coupled to an elemental analyser (NC2500 Carlo Erba) via a ConFloIII Interface. Due to the relatively wide range of C / N ratios of mosses, we used about 1.5 mg for each carbon stable isotope measurement (n = 400) and a replicate of 3 mg for each nitrogen stable isotope measurement (n = 326) and the analysis of elemental composition. The high weight needed for the nitrogen sample replicates prevented the measurement of $\delta^{15}N$ and thus the C / N for some samples. The calibration for carbon was performed using an urea standard and a δ^{13} C isotopic standard (IAEA CH-7). The nitrogen contents were calibrated against an acetanilide standard and the nitrogen isotopic composition with ammonium sulfate standard (IAEA N-1). The reliability of the method was checked with the NIST plant standard SRM 1547. The isotopic ratios are given in delta notation relative to Vienna Pee Dee Belemnite (VPDB) for δ^{13} C values and relative to air for δ^{15} N values. The reproducibility for replicate analyses is 0.2 % for carbon and nitrogen and 0.2 % for δ^{13} C and δ^{15} N values.

n-Alkane analyses were performed on a subset of 16 samples. We took material from the polygon complex P3 from the Khatanga region (72.149° N, 102.693° E), which comprises three successive low-centred polygons (Table S1) to exclude effects of environmental conditions at different locations. The moss samples were washed, identified and airdried. They were weighed (150–1000 mg dry weight) and samples were extracted with an accelerated solvent extractor (ASE) (Dionex, Sunnyvale; USA) using Cl₂Me:MeOH (10:1) at 5 bar and 75 °C. The extract was separated

Species	Abbreviation	<i>h</i> (mean plant position in relation to water level in centimetres)	Classification
Hylocomium splendens	Hyl_spl	16	xero-mesophilic
Tomentypnum nitens	Tom_nit	13	xero-mesophilic
Aulacomnium turgidum	Aul_tur	13	xero-mesophilic
Aulacomnium palustre	Aul_pal	13	xero-mesophilic
Hamatocaulis lapponicus	Ham_lap	3	meso-hygrophilic
Warnstorfia exannulata	War_exa	0*	meso-hygrophilic
Meesia triquetra	Mee_tri	-1	meso-hygrophilic
Drepanocladus revolvens	Dre_rev	-5	meso-hygrophilic
Scorpidium scorpioides	Sco_sco	-13	meso-hygrophilic
Calliergon giganteum	Cal_gig	-30	meso-hygrophilic

Table 1. Calculated mean plant position relative to the water level and the classification of the sampled brown-moss species.

* Samples of Warnstorfia exannulata were growing on a swinging mat at water level and mean plant position was thus set to zero.

into aliphatic hydrocarbon, aromatic hydrocarbon and nitrogen, and sulfur and oxygen compound fractions using medium-pressure liquid chromatography following Radke et al. (1980). A total of 5 µg of the quantification standard (5 α androstane, 1-ethylpyrene, 5α -androstan-17-one and erucic acid) was added. Gas chromatography (GC) of aliphatic hydrocarbon fractions was performed using a GC Agilent 6890 equipped with an Ultra-1 fused silica capillary column (Model Agilent 19091A-105, length 50 m, inner diameter 200 µm, film thickness 0.33 µm). Helium was used as a carrier gas at a constant flow rate of 1 mLmin^{-1} . The GC oven was heated from 40 °C (2 min hold time) to 300 °C (65 min hold time) at a rate of 5 °C min⁻¹. The samples were injected by means of splitless injection. Quantitative evaluation of data was done with ChemStation software. Additionally, we calculated the average chain length (ACL₂₁₋₃₃) assumed to represent a proxy for moisture (Andersson et al., 2011) and temperature (Bush and McInerney, 2015) with a comprehensible extended range from $n-C_{21}$ to $n-C_{33}$ and the proxy ratio $P_{\rm aq}$, which was developed as a proxy ratio to distinguish submerged or floating aquatic macrophytes from emergent and terrestrial plants (Ficken et al., 2000):

ACL_n =
$$\frac{\sum (n \times C_n)}{\sum C_n}$$
, $n = 21 - 33$,
 $P_{aq} = \frac{(C_{23} + C_{25})}{(C_{23} + C_{25} + C_{29} + C_{31})}$.

2.4 Statistical tests

We calculated the range, 0.25, 0.5, and 0.75 quantiles of δ^{13} C and δ^{15} N values, and the C/N ratios for all species (Table S2). Significant differences in the C/N ratios and δ^{13} C and δ^{15} N values among different groups were assessed with a *t* test. In addition, we performed linear regression between the obtained values for each individual species and surface height. All analyses were implemented in R version 3.2.0

(R Core Team, 2014). Furthermore, we performed a PCA with the percentage of n-alkane homologues (square-root transformed) in R using the "vegan" package version 2.0–10 (Oksanen et al., 2013).

3 Results

The overall ranges in C / N ratios of both groups have a broad overlap (xero-mesophilic: 22.5–67.9; meso-hygrophilic: 15.4–70.4). However, the medians of the C / N values of the xero-mesophilic species ranging from 47.6 to 52.9 (Fig. 2) are significantly higher than those of the meso-hygrophilic group, which range from 37.1 to 46.5 ($W = 18\,280$, $p \ll 0.001$). The C / N ratios show no intraspecific relations among individual species and water level (Fig. 3a), except for *Tomentypnum nitens* ($r^2 = 0.11$, p < 0.05).

The δ^{13} C values of the meso-hygrophilic group (-34.9 to -22.5%) cover the range of the xero-mesophilic group (-32.1 to -24.2%), which have a noticeably lower variance in δ^{13} C values (Fig. 2). The medians of the individual species in the xero-mesophilic group (range: -29.4 to -27.1%) are significantly different (W = 16232, p = 0.008) from those of the meso-hygrophilic group (range: -29.2 to -24.8%). δ^{13} C values of *Meesia triquetra*, *Drepanocladus revolvens*, and *Scorpidium scorpioides* (all belonging to the meso-hydrophilic group) are significantly positively related with the position of the water level (Fig. 3b), while no member of the xero-mesophilic group revealed such a relationship (Fig. 3a).

The ranges of δ^{15} N values of both groups are rather similar (Fig. 2). However, individual species medians of both groups are significantly different (t = -6.96, $p \ll 0.001$; xero-mesophilic group: -3.2 to -2.7%, meso-hygrophilic group: -2.9 to -0.1). Drepanocladus revolvens, Scorpidium scorpioides, and Calliergon giganteum, all belonging to the meso-hygrophilic group, exhibit a positive relationship be-



Figure 2. C / N ratios and δ^{13} C and δ^{15} N values of the 10 moss taxa studied. White box plots are the xero-mesophilic group and box plots shaded in blue are the meso-hygrophilic group. A *t* test was done to distinguish the signals between the two habitat groups.

tween the δ^{15} N values and position relative to water level (Fig. 3b).

n-Alkane distributions show the expected predominance of the odd chain length (Table 2). The absolute *n*-alkane concentrations (*n*-C₁₉ to *n*-C₃₃) range from 34 to 238 µg g⁻¹ of dry weight. The individual taxa show significant differences with respect to carbon number of the most abundant *n*alkane (*n*-C_{max}) forming unimodal distribution patterns. It is *n*-C₃₁ for *Tomentypnum nitens* (which is also the only species containing *n*-C₃₃ in detectable amounts), *n*-C₂₉ for *Aulacomnium turgidum*, *n*-C₂₇ for *Aulacomnium palustre*, *n*-C₂₇ and *n*-C₂₅ for the *Drepanocladus* taxa, and *n*-C_{max} = 25 for *Scorpidium scorpioides*. These differences in the *n*-alkane composition largely relate with the species-preferred mean plant position relative to water level (Fig. 4).

Evaluations of the *n*-alkane biomarker proxies, ACL_{21-33} and P_{aq} , also show a clear division between the xeromesophilic and the meso-hygrophilic species groups (Table 2), whereas intraspecific variations are rather small (with the exception of *Drepanocladus*). The xero-mesophilic group is notable for high averages of ACL_{21-33} (28.41) and P_{aq} (0.17) compared to low averages of ACL_{21-33} (25.61) and P_{aq} (0.87) for the meso-hygrophilic group.

The observed trend is also assumed in the biplot of the first two PCA axes, even though their explained variance is relatively low (16.9%) in the dataset (Fig. 5). The first axis separates xero-mesophilic from meso-hygrophilic taxa. *Aulacom*- *nium* taxa are located in the upper range and *Tomentypnum nitens* in the lower range of the second PCA axis, while no trend is observed within the meso-hygrophilic group along the second axis.

4 Discussion

4.1 C / N ratios

The C / N ratios of mosses from polygonal tundra in northern Siberia are relatively low compared with those obtained for mosses from Antarctic bogs that range between 80 and 100 (Björck et al., 1991) or from western Canada that range between 55 and 76 (Kuhry and Vitt, 1996). However, neither the taxa xeric nor mesic growing conditions were sampled in Antarctica and Canada. All investigated species are considered as ectohydric mosses, which receive nitrogen mostly from precipitation deposits (Ayres et al., 2006). Our results reveal that averaged C/N ratios for the xero-mesophilic moss group are higher than for the meso-hygrophilic group, probably reflecting the known difference between terrestrial and aquatic plants (Meyers and Ishiwatari, 1993). There are two possible impacts, which can influence the C / N ratio of these groups: (1) competition with vascular plants and (2) accessibility of nitrogen pools.



Figure 3. (a) Intraspecific relationships of the C / N ratio and the stable isotope values of carbon and nitrogen related to the water level of the xero-mesophilic moss group. Regression lines (red) are only plotted for significant datasets. (b) Intraspecific relationships of the C / N ratio and the stable isotope values of carbon and nitrogen related to the water level of the meso-hygrophilic moss group. Regression lines (red) are only plotted for significant datasets.

Species	<i>n</i> -C ₁₉	<i>n</i> -C ₂₁	<i>n</i> -C ₂₃	<i>n</i> -C ₂₅	<i>n</i> -C ₂₇	<i>n</i> -C ₂₉	<i>n</i> -C ₃₁	<i>n</i> -C ₃₃	ACL ₂₁₋₃₃	Paq
Xero-mesophilic habitat group										
Tom_nit (1)	0	2.0689	2.6374	6.0816	10.7454	17.7273	34.0742	0	28.86	0.14
Tom_nit (2)	0.5571	1.041	1.4493	2.6964	9.1008	8.4867	20.2818	7.2312	29.47	0.13
Tom_nit (3)	0.9936	1.456	2.175	2.8712	9.6118	7.2721	21.2309	4.9717	29.14	0.15
Aul_tur (1)	0	0.968	1.126	2.2001	8.9548	23.5267	3.5595	0	28.15	0.11
Aul_tur (2)	1.028	1.3468	1.2794	4.8797	18.7427	50.9291	5.5645	0	28.22	0.10
Aul_pal (1)	0	1.9298	2.5459	4.7646	20.5085	7.5826	7.979	0	27.35	0.32
Aul_pal (2)	0.7341	1.0767	1.5183	2.9728	12.3293	11.0909	4.9109	0	27.69	0.22
Meso-hygrophilic habitat group										
Dre_rev (1)	0.7868	1.5426	2.3659	43.833	19.0836	5.1551	3.5637	0	25.92	0.84
Dre_rev(2)	0.981	1.5227	2.7605	22.5638	14.5103	6.5356	7.5623	0	26.6	0.640
Dre_sp.	1.4696	1.9968	5.0986	29.6729	30.4582	9.9108	5.1317	0	26.38	0.70
$Sco_Sco(1)$	0	3.7612	11.7002	133.4207	29.0024	6.2023	1.9425	0	25.3	0.95
Sco_Sco (2)	0	3.8911	10.4693	93.7009	21.4601	5.8531	1.7706	0	25.29	0.93
$Sco_Sco(3)$	0.8856	2.7949	11.8988	134.378	24.2348	5.4035	2.4969	0	25.28	0.95
Sco_Sco (4)	1.6217	5.5813	11.705	101.7602	21.3126	6.9161	2.4522	0	25.26	0.92
$Sco_Sco(5)$	1.4083	3.8857	11.8966	121.5701	23.3167	4.315	1.7261	0	25.21	0.96
Sco_Sco (6)	1.345	2.6672	19.4796	170.3015	34.3255	6.8599	3.1732	0	25.28	0.95

- If moss plants invest in a high stem-to-leaf biomass ratio, which results in a high C / N ratio, they will increase their height and stability, and thus their competitive ability against vascular plants for light (Sveinbjörnsson and Oechel, 1992). Furthermore, the low N input by precipitation and a low N content of moss litter slows down the fungal and bacterial N mineralisation which increases the thickness of moss litter mats (Gornall et al., 2007; Turetsky, 2003). This in turn will increase the isolating function of moss mats, thus negatively affecting seed germination of vascular plants (Gornall et al., 2007).
- 2. Lower C/N ratios of meso-hygrophilic mosses may originate from higher amounts of dissolved nitrogen in polygon waters as a result of high net primary productivity, the presence of N₂-fixers such as cyanobacteria, and the exudations of zooplankton. Frahm (2001) assumes that loose epiphytic and endophytic symbiotic relationships between mosses and cyanobacteria are probably restricted to wetland taxa. Lindo et al. (2013) report such associations between brown mosses and cyanobacteria. Thus, the N supply is better for brown mosses preferring meso-hygrophilic than xero-mesophilic habitats, and the respective taxa accordingly have lower C/N ratios indicating habitat-specific variation in C / N ratios. We expected to also find intraspecific variations between C / N ratios and water level. The large variability in the C / N data may be a result of atmospheric conditions and organic matter degradation being the principal sources at xeric sites, whereas in mesic and wet sites

microbial symbionts play an important role in the C / N ratio. However, the signal-to-noise ratio is probably too low to give a meaningful result because only the average water level of each plot was recorded, but not that of each individual plant.

4.2 δ^{13} C values

With respect to bryophytes, most isotopic studies have hitherto been performed on *Sphagnum* (Markel et al., 2010; Ménot and Burns, 2001) while our study focuses on brown mosses – a major component in Siberian wetlands. The intraspecific variability for some meso-hygrophilic species (i.e. *Meesia triquetra, Drepanocladus revolvens, Scorpidium scorpioides*) show that the δ^{13} C signals are related to the hydrological conditions at the growing site of each individual, i.e. individuals growing at dry sites showed higher medial δ^{13} C values than those growing at wet sites. A difference among the two habitat groups is observed; they partly contradict the intraspecific findings in that some of the xeromesophilic species known to prefer dry rims such as *Hylocomium splendens* and *Tomentypnum nitens* have particularly low δ^{13} C medians.

The detected differences in moss δ^{13} C values, particularly of the meso-hygrophilic group, either reflect a source signal depending on water level or a physiological reaction of the plant related to water level (Bramley-Alves et al., 2014; Proctor et al., 1992). Mosses are typical C₃ plants (Farquhar et al., 1989; Rundel et al., 1979) characterized by a high CO₂ compensation point (Bain and Proctor, 1980; Dilks and Proc-



Figure 4. The relative amounts (%) of *n*-alkanes with an odd number of carbon atoms in selected brown mosses, grouped by species and ordered by their preferences with respect to their mean plant position relative to water level for xero-mesophilic (**a**, **b**) and meso-hygrophilic (**c**, **d**) mosses. Tom-nit: *Tomentypnum nitens*; Aul_tur: *Aulacomnium turgidum*; Aul_pal: *Aulacomnium palustre*; Dre_rev: *Drepanocladus revolvens*; Dre_sp: *Drepanocladus* sp.; Sco_sco: *Scorpidium scorpioides*.

tor, 1975; Salvucci and Bowes, 1981). The high availability of atmospheric CO₂ and elevated diffusion rates of CO₂ in air compared to water (O'Leary) result in typical terrestrial C₃ land plant δ^{13} C characteristics, because of a decreasing



Figure 5. PCA of *n*-alkanes separates the meso-hygrophilic group on the left side of axis 1 from the xero-mesophilic group on the right side. Along the second axis the *Aulacomniaceae* are distinguished from *Tomentypnum nitens* according to the distribution of long-chain *n*-alkanes with more or less than 30 carbon atoms.

cell water pressure in dry habitats which entails a strong discrimination rate against ¹³CO₂ induced by RuBisCO (Rice and Giles, 1996). With respect to the xero-mesophilic group, we observe an increase in discrimination against ¹³C from taxa preferring a low position relative to the water level (e.g. Aulacomnium taxa) than those preferring high positions (e.g. Hylocomium splendens). In contrast, if plant tissue is coated by a water film, the cell water pressure should reach an optimum, which is expected to result in a weaker discrimination rate against ¹³C by RuBisCO (Rice and Giles, 1996), because of a source restriction by the slower diffusion rate of CO₂ in water (Lloyd and Farquhar, 1994). A lower carbon isotope discrimination related to water saturation is observed for only three species out of six meso-hygrophilic mosses. However, this basic signal may be masked by variations in δ^{13} C values of different carbon sources, which are expected to be more influential for meso-hygrophilic mosses in watersaturated conditions. Ménot and Burns (2001) studied intraspecific variations for three Sphagnum species, which prefer three different habitat types (dry, meso, wet) along an elevational gradient, which was positively correlated with precipitation. They find a decline in discrimination against ¹³C with increasing wetness and, similar to our results, no relationship for species with a strong wetness preference. This is attributed to the variation in δ^{13} C from highly varying dissolved inorganic carbon (Proctor et al., 1992). Mosses potentially access ¹³C-depleted CO₂ that originates from oxidation of typically strongly ¹³C-depleted biogenic methane by methanotrophic microorganisms (Kip et al., 2010; Liebner et al., 2011; Raghoebarsing et al., 2005). Studies by Nichols et al. (2009) show that a higher water level at the peat surface is crucial for high methane-derived CO₂ release. Furthermore, symbiosis with methanotrophs enhances the moisture-related effect on the δ^{13} C signal of bryophytes. Endophytic microorganisms in hyalocytes of submerged Sphagnum (Raghoebarsing et al., 2005) or epiphytic microorganisms on submerged brown mosses (Liebner et al., 2011) are presumed to provide ¹³C-depleted CO₂ directly to the lamina cells of mosses. The studies of Ruttner (1947) and Bain and Proctor (1980) show that, in general, moss taxa are incapable of bicarbonate uptake. Hence, bicarbonate, known to be a carbon source for submerged vascular plants and algae (Herzschuh et al., 2010; Merz, 1992), can most probably be excluded as a carbon source for moss and thus as a factor influencing the δ^{13} C value. Moreover, the bicarbonate content in pond waters in northern Siberian landscapes is very low (Wetterich et al., 2008). Other sources of ¹³C-depleted CO₂ are surface run-off during spring flooding, rain events and decomposition processes in the pond (Leng and Marshall, 2004; Maberly et al., 2013). Yet we cannot fully eliminate the possibility that the measured bulk material was contaminated in parts with epiphytic or endophytic microorganisms. The overall isotopic composition would, however, likely be unaffected, as Ménot and Burns (2001) have shown that the δ^{13} C values of bulk organic material and alpha-cellulose of Sphagnum are very similar. Thus, the large ranges within several species of meso-hygrophilic habitats in arctic regions suggest that the existence of open water leads to more depleted δ^{13} C values, as well as measurements of the isotopic composition of methane when present, and microbial groups in the water and terrestrial litter should be possible. Finally, considering the relationship of selected brown mosses to mean plant position, the complex origin of plant-available carbon makes it difficult to interpret the δ^{13} C record, especially for meso-hygrophilic brown mosses as well as Sphagnum (Price et al., 1997; Ménot and Burns, 2001).

4.3 δ^{15} N isotopes

Like δ^{13} C, the interpretation of stable nitrogen isotope compositions of mosses is challenging because, again, source signals need to be separated from those originating from physiological isotopic discrimination processes. Our results yield relatively ¹⁵N-depleted δ^{15} N values for xero-mesophilic mosses growing preferentially on rims compared to mesohygrophilic mosses.

The terrestrial arctic systems are generally thought to be nitrogen-limited (Gordon et al., 2001; Kielland, 1997). On the rim sites, atmospheric deposition can be considered to be the most important source for nitrogen (Jonasson and Shaver, 1999) originating from fog, dew, precipitation and surface run-off (Sveinbjörnsson and Oechel, 1992). However, most of the nitrogen available to rim mosses originates from recycling of already ¹⁵N-depleted higher plant and moss litter

(Turetsky, 2003). The ectohydric morphology enables an efficient nutrient uptake across the entire moss plant surface via trapped water. In fact, the meso-hygrophilic group has a higher N content than the xero-mesophilic group (see section on C / N ratio). Inorganic nitrogen, and especially the high amounts of organic nitrogen provided by N-mineralisation in tundra soils (Kielland, 1995), are important for mosses growing on the rather dry sites such as the polygonal rims (Atkin, 1996).

Three of the investigated submerged or floating moss species show a significant positive relationship between water level and δ^{15} N values. These results are similar to those of Asada et al. (2005) who tested a relationship between δ^{15} N values of different Sphagnum species and their position relative to the groundwater level, which they assumed to originate from different nitrogen sources and different internal fractionating processes. We assume that the often-heavier nitrogen isotope composition of meso-hygrophilic brownmoss individuals originates from the high degree of symbiotic associations with aquatic atmospheric nitrogen-fixing autotrophic microorganisms such as Nostoc or Anabaena (Lindo et al., 2013), or methanotrophs (Vile et al., 2014). The high spatial degree with endo- or ectosymbiotic N₂-fixing microorganisms enables the direct uptake of their nitrogen products, which is similar to that of N₂ in air.

4.4 Patterns of *n*-alkane

Compared to vascular plants that are characterized by a thick leaf-wax layer, mosses produce only a small amount of *n*-alkanes (Baas et al., 2000; Ficken et al., 1998). Like previous studies on vascular plants (Aichner et al., 2010; Ficken et al., 2000; Meyers and Ishiwatari, 1993), our results generally reveal a differentiation between terrestrial taxa (i.e. xero-mesophilic group), characterized mainly by *n*-alkanes maximising at n-C₂₉ and n-C₃₁, and submerged living taxa (i.e. the meso-hygrophilic group) maximising at n-C₂₅ and n-C₂₇. Earlier investigations of Nott et al. (2000), Baas et al. (2000) and Bingham et al. (2010), who compared the *n*-alkane fingerprints of *Sphagnum* taxa growing along a hydrological gradient, agree with our results.

Huang et al. (2012a) and Ficken et al. (1998) used proxy ratios (ACL, P_{aq}) to divide moss taxa roughly by their moisture preferences. They calculate the ACL_{23–33} for samples of lichens and *Racomitrium lanuginosum*, which have rather similar hydrological requirements to our xero-mesophilic mosses. Despite the slightly narrower ACL_{23–33} range, their results show similarities to our xero-mesophilic group. A comparison between ACL_{21–33}, ratios of *Sphagnum* (plant position nearly at water level) of Huang et al. (2012b) and our brown mosses shows that the ACL_{21–33} ratios of *Sphagnum* species are rather lower. The intraspecific conclusion of Huang et al. (2012b) (wetter moisture conditions entail lower ACL₂₁₋₃₃) for *Sphagnum* is reflected by our measurements.

dry (ACL ₂₁₋₃₃		wet (ACL _{21–33}
= 29.1 - 27.5	< moisture condition <	= 26.4 - 25.2
Tom_nit < Aul_tur	< Aul_pal <	Dre_rev < Sco_sco

As we observed a clear difference in the ACL_{21-33} between the xero-mesophilic and the meso-hygrophilic group, we suggest that the inclusion of mid-chain *n*-alkanes (*n*- C_{21} to *n*- C_{25}) in the equation of ACL improves its value as a proxy for moisture conditions. Andersson et al. (2011) inferred ACL_{27-31} values of 29 for brown-moss peat from western Russia during wet phases, which is, however, poorly comparable to our results because they investigated total peat organic matter instead of pure moss material.

Ficken et al. (2000) proposed P_{aq} as a semi-quantitative proxy ratio for the differentiation of terrestrial and aquatic plants (< 0.1 terrestrial plants, 0.1–0.4 emergent macrophytes, 0.4–1 submerged or floating macrophytes). Our inferred P_{aq} results for the individual species agree with these assumptions. If we consider that the proxy ratio levels were created by vascular plants from a limited dataset of lakes in Kenya, and as we focus on non-vascular plants of the arctic, we chose other level terms.

emergent	plant types	submerged/floating
$(P_{oa} = 0.1 - 0.4)$	$<$ solution by $F_{aq} <$ (Ficken et al. 2000)	$(P_{aa} = 0.4-1)$
xero-mesophilic	adapted terms	meso-hygrophilic
mosses	for mosses	mosses
Tom nit < Aul tur	< Aul pal <	Dre rev $<$ Sco sco

Overall, our results do not support the inference of Nichols et al. (2006) that a hydrological classification is possible between *Sphagnum*- and non-*Sphagnum*-formed peat as the latter show wide variations between different habitats. The inferred broad P_{aq} range of *Drepanocladus* and between both *Aulacomnium* probably indicates that intraspecific variation is related to the individual's growing condition, which could provide the basis to develop P_{aq} as a proxy for water level when measured on taxonomically identified fossil plant material.

As with P_{aq} , *n*-alkanes seem to be species-specific given stable environmental parameters and are related to the species-specific moisture requirements, which are adapted to changing environmental conditions. Thus, our results confirm the conclusions of Bingham et al. (2010), Bush and McInerney (2015), and Nott et al. (2000) that the pattern of *n*-alkanes has the potential to become a valuable proxy for chemotaxonomic identification and moisture conditions. Scorpidium scorpioides, a species with a rather narrow preference range (i.e. it is limited to open-water conditions), shows low intraspecific variations. This matches the results for Sphagnum compiled by Bingham et al. (2010), which also show minor intraspecific variations. Aulacomnium, in contrast, which grows in a rather wide range of moisture conditions, shows strong variations in its *n*-alkane spectra: whether this is a function of the individual's growing conditions, however, needs to be investigated in a more extensive study.

5 Conclusions

The habitat and intraspecific isotopic and chemical patterns of 10 brown-moss species detected along small-scale hydrological gradients in Siberian polygonal tundra were studied.

The observed higher C/N ratios of xero-mesophilic mosses compared to those of the meso-hygrophilic mosses originate from the different environmental requirements when living emergent (i.e. investment in a higher stability resulting in high C/N ratios) as opposed to submerged. Furthermore, the latter group may also gain a better nitrogen supply through microbial symbioses.

With respect to the isotopic source pools, the mesohygrophilic species have greater access than xero-mesophilic species, which is seen in their large ranges. The approximate habitat-specific division of δ^{13} C values as a result of discrimination by RuBisCO under different hydrological regimes is overturned by the influence of different sources and cannot provide a clear distinction from a single measurement of either habitat type. For species growing near the water level, no intraspecific relationship with water level was observed, probably as a result of the parallel impact of processes causing opposing δ^{13} C trends.

Our analyses reveal that, compared with xero-mesophilic mosses, meso-hygrophilic mosses are characterized by enriched δ^{15} N values probably originating from microbial symbioses. Both carbon and nitrogen isotopic ratios seem to be valuable proxies to differentiate between taxa preferring the polygon rim or pond. Moreover, with respect to meso-hygrophilic mosses, the detected positive relations between intraspecific variations and the individuals' relative growing position could allow even more semi-quantitative information about water-level changes to be inferred.

The *n*-alkane patterns of brown mosses (limited 16 individuals belonging to 5 species) indicate that they are species-specific and thus have the potential to be developed as a chemotaxonomic proxy. The applicability of proxy ratios (ACL and P_{aq}) could be attested for arctic mosses after adjustments of the levels.

Overall, our study indicates that C/N, isotopic and *n*-alkane analyses of brown-moss material has a high environmental indicator potential, particularly if species-specific material instead of bulk material is analysed.

Data availability. The Supplement related to this article is available online at PANGAEA, the Data Publisher for Earth & Environmental Science; https://doi.org/10.1594/PANGAEA.873765 (Zibulski et al., 2017).

The Supplement related to this article is available online at doi:10.5194/bg-14-1617-2017-supplement.

Competing interests. The authors declare that they have no conflict of interest.

The article processing charges for this open-access publication were covered by a Research Centre of the Helmholtz Association.

Edited by: Silvio Pantoja Reviewed by: P. A. Meyers and one anonymous referee

References

- Aerts, R., Verhoeven, J. T. A., and Whigham, D. F.: Plant-mediated controls on nutrient cycling in temperate fens and bogs, Ecology, 80, 2170–2181, 1999.
- Aichner, B., Herzschuh, U., and Wilkes, H.: Influence of aquatic macrophytes on the stable carbon isotopic signatures of sedimentary organic matter in lakes on the Tibetan Plateau, Org. Geochem., 41, 706–718, doi:10.1016/j.orggeochem.2010.02.002, 2010.
- Andersson, R. A., Kuhry, P., Meyers, P., Zebühr, Y., Crill, P., and Mörth, M.: Impacts of paleohydrological changes on n-alkane biomarker compositions of a Holocene peat sequence in the eastern European Russian Arctic, Org. Geochem., 42, 1065–1075, doi:10.1016/j.orggeochem.2011.06.020, 2011.
- Asada, T., Warner, B. G., and Aravena, R.: Nitrogen isotope signature variability in plant species from open peatland, Aquat. Bot., 82, 297–307, doi:10.1016/j.aquabot.2005.05.005, 2005.
- Atkin, O. K.: Reassessing the nitrogen relations of Arctic plants: a mini-review, Plant Cell Environ., 19, 695–704, doi:10.1111/j.1365-3040.1996.tb00404.x, 1996.
- Ayres, E., van der Wal, R., Sommerkorn, M., and Bardgett, R. D.: Direct uptake of soil nitrogen by mosses, Biol. Lett., 2, 286–288, doi:10.1098/rsbl.2006.0455, 2006.
- Baas, M., Pancost, R., van Geel, B., and Sinninghe Damsté, J. S.: A comparative study of lipids in *Sphagnum* species, Org. Geochem., 31, 535–541, doi:10.1016/S0146-6380(00)00037-1, 2000.
- Bain, J. T. and Proctor, M. C. F.: The requirement of aquatic bryophytes for free CO₂ as an inorganic carbon source: Some experimental Evidence, New Phytol., 86, 393–400, doi:10.1111/j.1469-8137.1980.tb01680.x, 1980.
- Beringer, J., Lynch, A. H., Chapin, F. S., Mack, M., and Bonan, G. B.: The representation of Arctic soils in the Land Surface Model: The importance of mosses, J. Climate, 14, 3324–3335, doi:10.1175/1520-0442(2001)014<3324:TROASI>2.0.CO;2, 2001.
- Bingham, E. M., McClymont, E. L., Väliranta, M., Mauquoy, D., Roberts, Z., Chambers, F. M., Pancost, R. D., and Evershed, R.
 P.: Conservative composition of *n*-alkane biomarkers in *Sphagnum* species: implications for palaeoclimate reconstruction in ombrotrophic peat bogs, Org. Geochem., 41, 214–220, 2010.
- Birks, H. J. B.: Quaternary bryophyte palaeoecology, in: Bryophyte Ecology, edited by: Smith, A. J. E., 473–490, Springer, the Netherlands, available at: http://link.springer.com/chapter/10. 1007/978-94-009-5891-3_13 (last access: 2 March 2014), 1982.
- Björck, S., Malmer, N., Hjort, C., Sandgren, P., Ingólfsson, Ó., Wallén, B., Smith, R. I. L., and Jónsson, B. L.: Stratigraphic and paleoclimatic studies of a 5500-year-old moss bank on

Elephant Island, Antarctica, Arct. Alp. Res., 23, 361–374, doi:10.2307/1551679, 1991.

- Blok, D., Heijmans, M. M. P. D., Schaepman-Strub, G., Ruijven, J. van, Parmentier, F. J. W., Maximov, T. C., and Berendse, F.: The cooling capacity of mosses: Controls on water and energy fluxes in a Siberian tundra site, Ecosystems, 14, 1055–1065, doi:10.1007/s10021-011-9463-5, 2011.
- Boike, J., Kattenstroth, B., Abramova, K., Bornemann, N., Chetverova, A., Fedorova, I., Fröb, K., Grigoriev, M., Grüber, M., Kutzbach, L., Langer, M., Minke, M., Muster, S., Piel, K., Pfeiffer, E.-M., Stoof, G., Westermann, S., Wischnewski, K., Wille, C., and Hubberten, H.-W.: Baseline characteristics of climate, permafrost and land cover from a new permafrost observatory in the Lena River Delta, Siberia (1998–2011), Biogeosciences, 10, 2105–2128, doi:10.5194/bg-10-2105-2013, 2013.
- Bramley-Alves, J., Wanek, W., French, K., and Robinson, S. A.: Moss δ^{13} C: an accurate proxy for past water environments in polar regions, available at: http://www.researchgate.net/profile/ Jessica_Bramley-Alves/publication/270164091_Moss_13C_ an_accurate_proxy_for_past_water_environments_in_polar_ regions/links/54ea77150cf27a6de1146b82.pdf (last access: 3 September 2015), 2014.
- Bush, R. T. and McInerney, F. A.: Leaf wax *n*-alkane distributions in and across modern plants: implications for paleoecology and chemotaxonomy, Geochim. Cosmochim. Ac., 117, 161–179, 2013.
- Bush, R. T. and McInerney, F. A.: Influence of temperature and C₄ abundance on n-alkane chain length distributions across the central USA, Org. Geochem., 117, 161–179, doi:10.1016/j.orggeochem.2014.12.003, 2015.
- Chambers, F. M. and Charman, D. J.: Holocene environmental change: contributions from the peatland archive, Holocene, 14, 1–6, 2004.
- Chanway, C. P., Anand, R., and Yang, H.: Nitrogen Fixation Outside and Inside Plant Tissues, available at: http://cdn.intechopen.com/ pdfs-wm/46070.pdf (last access: 5 May 2015), 2014.
- Dilks, T. J. K. and Proctor, M. C. F.: Comparative experiments on temperature responses of bryophytes: assimilation, respiration and freezing damage, J. Bryol., 8, 317–336, 1975.
- Farquhar, G. D., Ehleringer, J. R., and Hubick, K. T.: Carbon isotope discrimination and photosynthesis, Annu. Rev. Plant Biol., 40, 503–537, 1989.
- Feakins, S. J., Peters, T., Wu, M. S., Shenkin, A., Salinas, N., Girardin, C. A. J., Bentley, L. P., Blonder, B., Enquist, B. J., Martin, R. E., Asner, G. P., and Malhi, Y.: Production of leaf wax n-alkanes across a tropical forest elevation transect, Org. Geochem., 100, 89–100, doi:10.1016/j.orggeochem.2016.07.004, 2016.
- Ficken, K. J., Barber, K. E., and Eglinton, G.: Lipid biomarker, δ^{13} C and plant macrofossil stratigraphy of a Scottish montane peat bog over the last two millennia, Org. Geochem., 28, 217–237, 1998.
- Ficken, K. J., Li, B., Swain, D. L., and Eglinton, G.: An *n*-alkane proxy for the sedimentary input of submerged/floating freshwater aquatic macrophytes, Org. Geochem., 31, 745–749, doi:10.1016/S0146-6380(00)00081-4, 2000.
- Frahm, J. P.: Biologie der Moose, Spektrum Akademischer Verlag, 2001.
- Frahm, J.-P. and Frey, W.: Moosflora, 4., Neubearbeitete und erweiterte Auflage, UTB, Stuttgart, 2004.

- Galimov, E. M.: Carbon isotope composition of Antarctic plants, Geochim. Cosmochim. Ac., 64, 1737–1739, doi:10.1016/S0016-7037(99)00328-2, 2000.
- Gordon, C., Wynn, J. M., and Woodin, S. J.: Impacts of increased nitrogen supply on high Arctic heath: the importance of bryophytes and phosphorus availability, New Phytol., 149, 461– 471, doi:10.1046/j.1469-8137.2001.00053.x, 2001.
- Gornall, J. L., Jónsdóttir, I. S., Woodin, S. J., and van der Wal, R.: Arctic mosses govern below-ground environment and ecosystem processes, Oecologia, 153, 931–941, 2007.
- Gornall, J. L., Woodin, S. J., Jónsdóttir, I. S., and van der Wal, R.: Balancing positive and negative plant interactions: how mosses structure vascular plant communities, Oecologia, 166, 769–782, 2011.
- Harmens, H., Norris, D. A., Cooper, D. M., Mills, G., Steinnes, E., Kubin, E., Thöni, L., Aboal, J. R., Alber, R., Carballeira, A., Coşkun, M., De Temmerman, L., Frolova, M., González-Miqueo, L., Jeran, Z., Leblond, S., Liiv, S., Maňkovská, B., Pesch, R., Poikolainen, J., Rühling, Å., Santamaria, J. M., Simonèiè, P., Schröder, W., Suchara, I., Yurukova, L., and Zechmeister, H. G.: Nitrogen concentrations in mosses indicate the spatial distribution of atmospheric nitrogen deposition in Europe, Environ. Pollut., 159, 2852–2860, doi:10.1016/j.envpol.2011.04.041, 2011.
- Herzschuh, U., Mischke, S., Meyer, H., Plessen, B., and Zhang, C.: Using variations in the stable carbon isotope composition of macrophyte remains to quantify nutrient dynamics in lakes, J. Paleolimnol., 43, 739–750, doi:10.1007/s10933-009-9365-0, 2010.
- Huang, X., Xue, J., Zhang, J., Qin, Y., Meyers, P. A., and Wang, H.: Long chain n -alkanes and their carbon isotopes in lichen species from western Hubei Province: implication for geological records, Front. Earth Sci., 6, 95–100, doi:10.1007/s11707-012-0300-8, 2012a.
- Huang, X., Xue, J., Zhang, J., Qin, Y., Meyers, P. A., and Wang, H.: Effect of different wetness conditions on Sphagnum lipid composition in the Erxianyan peatland, central China, Org. Geochem., 44, 1–7, 2012b.
- Jonasson, S. and Shaver, G. R.: Within-stand nutrient cycling in arctic and boreal wetlands, Ecology, 80, 2139–2150, doi:10.1890/0012-9658(1999)080[2139:WSNCIA]2.0.CO;2, 1999.
- Kielland, K.: Landscape patterns of free amino acids in arctic tundra soils, Biogeochemistry, 31, 85–98, doi:10.1007/BF00000940, 1995.
- Kielland, K.: Role of free amino acids in the nitrogen economy of arctic cryptogams, Ecoscience, 4, 75–79, 1997.
- Kip, N., van Winden, J. F., Pan, Y., Bodrossy, L., Reichart, G.-J., Smolders, A. J. P., Jetten, M. S. M., Damsté, J. S. S., and Op den Camp, H. J. M.: Global prevalence of methane oxidation by symbiotic bacteria in peat-moss ecosystems, Nat. Geosci., 3, 617–621, doi:10.1038/ngeo939, 2010.
- Kuhry, P. and Vitt, D. H.: Fossil carbon/nitrogen ratios as a measure of peat decomposition, Ecology, 77, 271–275, doi:10.2307/2265676, 1996.
- Lee, Y. I., Lim, H. S., and Yoon, H. I.: Carbon and nitrogen isotope composition of vegetation on King George Island, maritime Antarctic, Polar Biol., 32, 1607–1615, doi:10.1007/s00300-009-0659-5, 2009.

- Leng, M. J. and Marshall, J. D.: Palaeoclimate interpretation of stable isotope data from lake sediment archives, Quaternary Sci. Rev., 23, 811–831, 2004.
- Liebner, S., Zeyer, J., Wagner, D., Schubert, C., Pfeiffer, E.-M., and Knoblauch, C.: Methane oxidation associated with submerged brown mosses reduces methane emissions from Siberian polygonal tundra, J. Ecol., 99, 914–922, doi:10.1111/j.1365-2745.2011.01823.x, 2011.
- Lindo, Z., Nilsson, M.-C., and Gundale, M. J.: Bryophytecyanobacteria associations as regulators of the northern latitude carbon balance in response to global change, Glob. Change Biol., 19, 2022–2035, 2013.
- Liu, X.-Y., Xiao, H.-Y., Liu, C.-Q., Li, Y.-Y., and Xiao, H.-W.: Stable carbon and nitrogen isotopes of the moss *Hap-locladium microphyllum* in an urban and a background area (SW China): The role of environmental conditions and atmospheric nitrogen deposition, Atmos. Environ., 42, 5413–5423, doi:10.1016/j.atmosenv.2008.02.038, 2008.
- Lloyd, J. and Farquhar, G. D.: ¹³C discrimination during CO₂ assimilation by the terrestrial biosphere, Oecologia, 99, 201–215, 1994.
- Loader, N. J., McCarroll, D., van der Knaap, W. O., Robertson, I., and Gagen, M.: Characterizing carbon isotopic variability in *Sphagnum*, The Holocene, 17, 403–410, doi:10.1177/0959683607076474, 2007.
- Lobin, W., Fischer, E., Frahm, J.-P., Frey, W., and Gams, H.: Kleine Kryptogamenflora, Bd. 4, Die Moospflanzen und Farnpflanzen Europas, 6th Edn., Spektrum Akademischer Verlag, TS21 Gustav-Fischer-Verlag, Stuttgart, 1995.
- Maberly, S. C., Barker, P. A., Stott, A. W., and De Ville, M. M.: Catchment productivity controls CO₂ emissions from lakes, Nature Climate Change, 3, 391–394, 2013.
- Malmer, N., Svensson, B. M., and Wallén, B.: Interactions between *Sphagnum* mosses and field layer vascular plants in the development of peat-forming systems, Folia Geobot. Phytotx., 29, 483– 496, 1994.
- Markel, E. R., Booth, R. K., and Qin, Y.: Testate amoebae and δ^{13} C of Sphagnum as surface-moisture proxies in Alaskan peatlands, Holocene, 20, 463–475, doi:10.1177/0959683609354303, 2010.
- Matveev, I. A. (Ed.): Agricultural Atlas of the Republic Sakha (Yakutia), Nauka, Moscow, 1989.
- McFadden, J. P., Eugster, W., and Chapin, F. S.: A regional study of the controls on water vapor and CO₂ exchange in arctic tundra, Ecology, 84, 2762–2776, doi:10.1890/01-0444, 2003.
- Ménot, G. and Burns, S. J.: Carbon isotopes in ombrogenic peat bog plants as climatic indicators: calibration from an altitudinal transect in Switzerland, Org. Geochem., 32, 233–245, doi:10.1016/S0146-6380(00)00170-4, 2001.
- Merz, D. M. U. E.: The biology of carbonate precipitation by cyanobacteria, Facies, 26, 81–101, doi:10.1007/BF02539795, 1992.
- Meyers, P. A. and Ishiwatari, R.: Lacustrine organic geochemistry – an overview of indicators of organic matter sources and diagenesis in lake sediments, Org. Geochem., 20, 867–900, doi:10.1016/0146-6380(93)90100-P, 1993.
- Michelsen, A., Schmidt, I. K., Jonasson, S., Quarmby, C., and Sleep, D.: Leaf ¹⁵N abundance of subarctic plants provides field evidence that ericoid, ectomycorrhizal and non- and arbuscular

mycorrhizal species access different sources of soil nitrogen, Oecologia, 105, 53–63, doi:10.1007/BF00328791, 1996.

- Michelsen, A., Quarmby, C., Sleep, D., and Jonasson, S.: Vascular plant ¹⁵N natural abundance in heath and forest tundra ecosystems is closely related with presence and type of mycorrhizal fungi in roots, Oecologia, 115, 406–418, 1998.
- Mutke, J. and Barthlott, W.: Patterns of vascular plant diversity at continental to global scales, Biol. Skrif., 55, 521–531, 2005.
- Nebel, M. and Philippi, G.: Die Moose Baden-Württembergs, Ulmer, Stuttgart, Germany, 2005.
- Nichols, J. E., Booth, R. K., Jackson, S. T., Pendall, E. G., and Huang, Y.: Paleohydrologic reconstruction based on n-alkane distributions in ombrotrophic peat, Org. Geochem., 37, 1505– 1513, doi:10.1016/j.orggeochem.2006.06.020, 2006.
- Nichols, J. E., Walcott, M., Bradley, R., Pilcher, J., and Huang, Y.: Quantitative assessment of precipitation seasonality and summer surface wetness using ombrotrophic sediments from an Arctic Norwegian peatland, Quaternary Res., 72, 443–451, 2009.
- Nott, C. J., Xie, S., Avsejs, L. A., Maddy, D., Chambers, F. M., and Evershed, R. P.: *n*-Alkane distributions in ombrotrophic mires as indicators of vegetation change related to climatic variation, Org. Geochem., 31, 231–235, doi:10.1016/S0146-6380(99)00153-9, 2000.
- Oksanen, J., Blanchet, F. G., Kindt, R., Legendre, P., Minchin, P. R., O'Hara, R. B., Simpson, G. L., Solymos, P., Stevens, M. H. H., and Wagner, H.: Package "vegan", Community Ecol. Package Version, 2.0–10, available at: http://www.R-project.org (last access: 22 October 2015), 2013.
- O'Leary, M. H.: Carbon isotopes in photosynthesis, Bioscience, 38, 328–336, 1988.
- Pancost, R. D., van Geel, B., Baas, M., and Damsté, J. S. S.: ?13C values and radiocarbon dates of microbial biomarkers as tracers for carbon recycling in peat deposits, Geology, 28, 663–666, 2000
- Poikolainen, J., Piispanen, J., Karhu, J., and Kubin, E.: Long-term changes in nitrogen deposition in Finland (1990–2006) monitored using the moss *Hylocomium splendens*, Environ. Pollut., 157, 3091–3097, doi:10.1016/j.envpol.2009.05.018, 2009.
- Prahl, F. G., Bennett, J. T., and Carpenter, R.: The early diagenesis of aliphatic hydrocarbons and organic matter in sedimentary particulates from Dabob Bay, Washington, Geochim. Cosmochim. Ac., 44, 1967–1976, 1980.
- Price, G. D., McKenzie, J. E., Pilcher, J. R., and Hoper, S. T.: Carbon-isotope variation in *Sphagnum* from hummock-hollow complexes: implications for Holocene climate reconstruction, Holocene, 7, 229–233, doi:10.1177/095968369700700211, 1997.
- Proctor, M. C. F., Raven, J. A., and Rice, S. K.: Stable carbon isotope discrimination measurements in *Sphagnum* and other bryophytes: physiological and ecological implications, J. Bryol., 17, 193–202, 1992.
- Radke, M., Willsch, H., and Welte, D. H.: Preparative hydrocarbon group type determination by automated medium pressure liquid chromatography, Anal. Chem., 52, 406–411, 1980.
- Raghoebarsing, A. A., Smolders, A. J. P., Schmid, M. C., Rijpstra, W. I. C., Wolters-Arts, M., Derksen, J., Jetten, M. S. M., Schouten, S., Sinninghe Damsté, J. S., Lamers, L. P. M., Roelofs, J. G. M., Op den Camp, H. J. M., and Strous, M.: Methanotrophic

symbionts provide carbon for photosynthesis in peat bogs, Nature, 436, 1153–1156, doi:10.1038/nature03802, 2005.

- R Core Team: R: A Language and Environment for Statistical Computing, R Foundation for Statistical Computing, Vienna, Austria, available at: http://www.R-project.org (last access: 15 November 2014), 2014.
- Rice, S. K. and Giles, L.: The influence of water content and leaf anatomy on carbon isotope discrimination and photosynthesis in *Sphagnum*, Plant Cell Environ., 19, 118–124, doi:10.1111/j.1365-3040.1996.tb00233.x, 1996.
- Rivas-Martinez, S. and Rivas-Saenz, S.: Worldwide Bioclimatic Classification System, 1996–2009, Phytosciological Res. Cent. Spain, available at: http://www.globalbioclimatics.org/default. htm (last access: 25 October 2012), 2007.
- Rundel, P. W., Stichler, W., Zander, R. H., and Ziegler, H.: Carbon and hydrogen isotope ratios of bryophytes from arid and humid regions, Oecologia, 44, 91–94, doi:10.1007/BF00346404, 1979.
- Ruttner, F.: Zur Frage der Karbonatassimilation der Wasserpflanzen, Plant Syst. Evol., 94, 265–294, 1947.
- Sachse, D., Radke, J., and Gleixner, G.: δD values of individual *n*-alkanes from terrestrial plants along a climatic gradient – Implications for the sedimentary biomarker record, Org. Geochem., 37, 469–483, doi:10.1016/j.orggeochem.2005.12.003, 2006.
- Salvucci, M. E. and Bowes, G.: Induction of reduced photorespiratory activity in submersed and amphibious aquatic macrophytes, Plant Physiol., 67, 335–340, 1981.
- Schellekens, J. and Buurman, P.: *n*-Alkane distributions as palaeoclimatic proxies in ombrotrophic peat: The role of decomposition and dominant vegetation, Geoderma, 164, 112–121, doi:10.1016/j.geoderma.2011.05.012, 2011.
- Skrzypek, G., Kałużny, A., Wojtuń, B., and Jędrysek, M.-O.: The carbon stable isotopic composition of mosses: A record of temperature variation, Org. Geochem., 38, 1770–1781, doi:10.1016/j.orggeochem.2007.05.002, 2007.
- Smith, A. J. E.: The Moss Flora of Britain and Ireland, Cambridge University Press, Cambridge, 1978.
- Smith, B. N. and Epstein, S.: Two categories of ¹³C / ¹²C ratios for higher plants, Plant Physiol., 47, 380–384, doi:10.1104/pp.47.3.380, 1971.
- Sveinbjörnsson, B. and Oechel, W. C.: Controls on growth and productivity of bryophytes: environmental limitations under current and anticipated conditions, in: Bryophytes and lichens in a changing environment, edited by: Bates, J. W. and Farmer, A. M., 77–102, Oxford University Press, Oxford, 1992.
- Tarnocai, C., Canadell, J. G., Schuur, E. A. G., Kuhry, P., Mazhitova, G., and Zimov, S.: Soil organic carbon pools in the northern circumpolar permafrost region, Global Biogeochem. Cy., 23, GB2023, doi:10.1029/2008GB003327, 2009.
- Turetsky, M. R.: The role of bryophytes in carbon and nitrogen cycling, Bryologist, 106, 395–409, doi:10.1639/05, 2003.
- Turetsky, M. R., Bond-Lamberty, B., Euskirchen, E., Talbot, J., Frolking, S., McGuire, A. D., and Tuittila, E.-S.: The resilience and functional role of moss in boreal and arctic ecosystems, New Phytol., 196, 49–67, doi:10.1111/j.1469-8137.2012.04254.x, 2012.
- Vile, M. A., Wieder, R. K., Živković, T., Scott, K. D., Vitt, D. H., Hartsock, J. A., Iosue, C. L., Quinn, J. C., Petix, M., Fillingim, H. M., Popma, J. M. A., Dynarski, K. A., Jackman, T. R., Albright, C. M., and Wykoff, D. D.: N₂-fixation by methanotrophs

sustains carbon and nitrogen accumulation in pristine peatlands, Biogeochemistry, 121, 317–328, doi:10.1007/s10533-014-0019-6, 2014.

- Waite, M. and Sack, L.: Shifts in bryophyte carbon isotope ratio across an elevation × soil age matrix on Mauna Loa, Hawaii: do bryophytes behave like vascular plants?, Oecologia, 166, 11–22, doi:10.1007/s00442-010-1903-y, 2011.
- Wasley, J., Robinson, S. A., Lovelock, C. E., and Popp, M.: Some like it wet – biological characteristics underpinning tolerance of extreme water stress events in Antarctic bryophytes, Funct. Plant Biol., 33, 443–455, 2006.
- Wetterich, S., Schirrmeister, L., Meyer, H., Viehberg, F. A., and Mackensen, A.: Arctic freshwater ostracods from modern periglacial environments in the Lena River Delta (Siberian Arctic, Russia): geochemical applications for palaeoenvironmental reconstructions, J. Paleolimnol., 39, 427–449, doi:10.1007/s10933-007-9122-1, 2008.
- Zhou, Y., Grice, K., Stuart-Williams, H., Farquhar, G. D., Hocart, C. H., Lu, H., and Liu, W.: Biosynthetic origin of the saw-toothed profile in δ^{13} C and δ^{2} H of n-alkanes and systematic isotopic differences between n-, iso- and anteiso-alkanes in leaf waxes of land plants, Phytochemistry, 71, 388–403, doi:10.1016/j.phytochem.2009.11.009, 2010.
- Zibulski, R., Herzschuh, U., and Pestryakova, L. A.: Vegetation patterns along micro-relief and vegetation type transects in polygonal landscapes of the Siberian Arctic, J. Veg. Sci., 27, 377–386, doi:10.1111/jvs.12356, 2016.
- Zibulski, R., Wesener, F., Wilkes, H., Plessen, B., Pestryakova, L. A., and Herzschuh, U.: Stable isotope (δ^{13} C, δ^{15} N), and n-alkane patterns of bryophytes along hydrological gradients of low-centred polygon of the Siberian Arctic, doi:10.1594/PANGAEA.873765, 2017.